
Chapter 2 The Highland Border Complex

[Excursion 1 Arbroath](#)

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In the Fife and Angus area, that part lying within the Midland Valley of Scotland is underlain by Upper Palaeozoic rocks belonging to the Devonian and Carboniferous Systems, part of the Midland Valley Block of Curry *et al.* (1984). Beyond the Midland Valley to the north-west lie the Scottish Highlands, underlain by much older Dalradian rocks of the Precambrian, the Dalradian Block of Curry *et al.* (1984). These two major areas are separated by the Highland Boundary Fault, not a simple single fault plane but rather a series of anastomosing faults between which occur fault-bounded wedges of rock seldom more than a kilometre wide. These wedges are largely composed of the rocks of the Highland Border Complex which were, until recently, thought to have an affinity with the adjacent Dalradian (e.g. Harris and Pitcher 1975).

More recent work indicates that the Highland Boundary Fault is a major boundary separating 'terranes' originally far from their present position and brought together by means of very large horizontal or transform displacements. Caught up within this major tectonic boundary are the wedges or 'slivers' (Curry *et al.* 1984) assigned to the Highland Border Complex.

Within the area covered by the guide book, fossils have been obtained from rocks of the Highland Border Complex at three levels: in Glen Esk (1) in the Margie Limestone — Upper Ordovician (Burton *et al.* 1984) and (2) in 'greenstones' — Lower Ordovician (Downie *et al.* 1971, p. 24), and at Stone-haven in cherty shales and interbedded greenstones/volcanics — probably Middle Ordovician (Curry *et al.* 1984).

Modern stratigraphical practice (Curry *et al.* 1984) suggests that Barrow's (1901) terms Jasper and Greenrock Series and Margie Series as applied to the North Esk and Stonehaven sections should be dropped and the informal terms 'rock assemblage' used until such time as more closely defined terms can be set up. The terms assemblages 1, 3 and 4 allow comparison with the better documented Aberfoyle area and are used below with the older Barrow terms in brackets for convenience. A wide variety of rock types are found in Rock Assemblage 1 including serpentinite; Rock Assemblage 3 contains black shales, cherts and volcanics, e.g. pillow lavas (Barrow's Jasper and Greenrock Series), and finally Rock Assemblage 4 comprises breccias, conglomerates and arenites and, in Glen Esk, limestone (Barrow's Margie Series).

Of the Highland Border Complex the following can be said (Curry *et al.* 1984):

1. The complex is younger than both the Tay Nappe and the peak of regional metamorphism in the Dalradian see (Table 2).
2. The complex is made up of a wide range of lithologies brought together in a series of fault slivers, and ranging in age through most of the Ordovician and according to Rogers *et al.* (1989) Lower Cambrian.
3. Omitting for the moment the Cambrian, the Highland Border Complex has a varied stratigraphical history comprising an older pre-Arenig serpentinite complex, overlain unconformably by carbonates (in the Aberfoyle area), conglomerates, black shales, cherts and volcanics of Lower and Middle Ordovician age. These in turn are overlain unconformably by Upper Ordovician arenites, shales, limestones and conglomerates.
4. The sequence occurs along the important boundary between a poorly known Midland Valley basement, a possible source of much of the Highland Border Complex sediments, and the Dalradian rocks of the Highlands.

The rocks of the Highland Border Complex may well have been formed in an oceanic marginal basin with the sediment source on the south-east side. Of the adjacent Dalradian rocks there is little or no evidence until near the end of the Lower Palaeozoic, a time of Dalradian uplift and retrogressive metamorphism and the formation of the Highland Border down-bend see (Table 2), and it seems likely that only at this time did the rocks of the Dalradian and the Highland Border Complex come together as a result of major transcurrent faulting (Harte *et al.* 1984, p. 162). Such major transcurrent

faulting may also have been responsible for the introduction of the Cambrian Leny Limestone to the Callander area further west.

References

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<i>Time (My)</i>	<i>Structures</i>	<i>Metamorphism</i>	<i>Uplift</i>	<i>Igneous Events</i>
pre 590	D1/D2 Tay and Tarfside Nappe	Low grade		
550				
500	D3	Peak metamorphism	Local in Upper Glen Esk	
450	D4 Highland Border Downbend	Retrogressive Metamorphism	Regional Uplift	
400				Post-tectonic Granites, e.g. Comrie

(Table 2) Dalradian evolution (After Rogers *et al.* 1989 and Harte *et al.* 1984).